

DE LA RECHERCHE À L'INDUSTRIE

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Laboratoire d'Hydrologie et
de Géochimie de Strasbourg

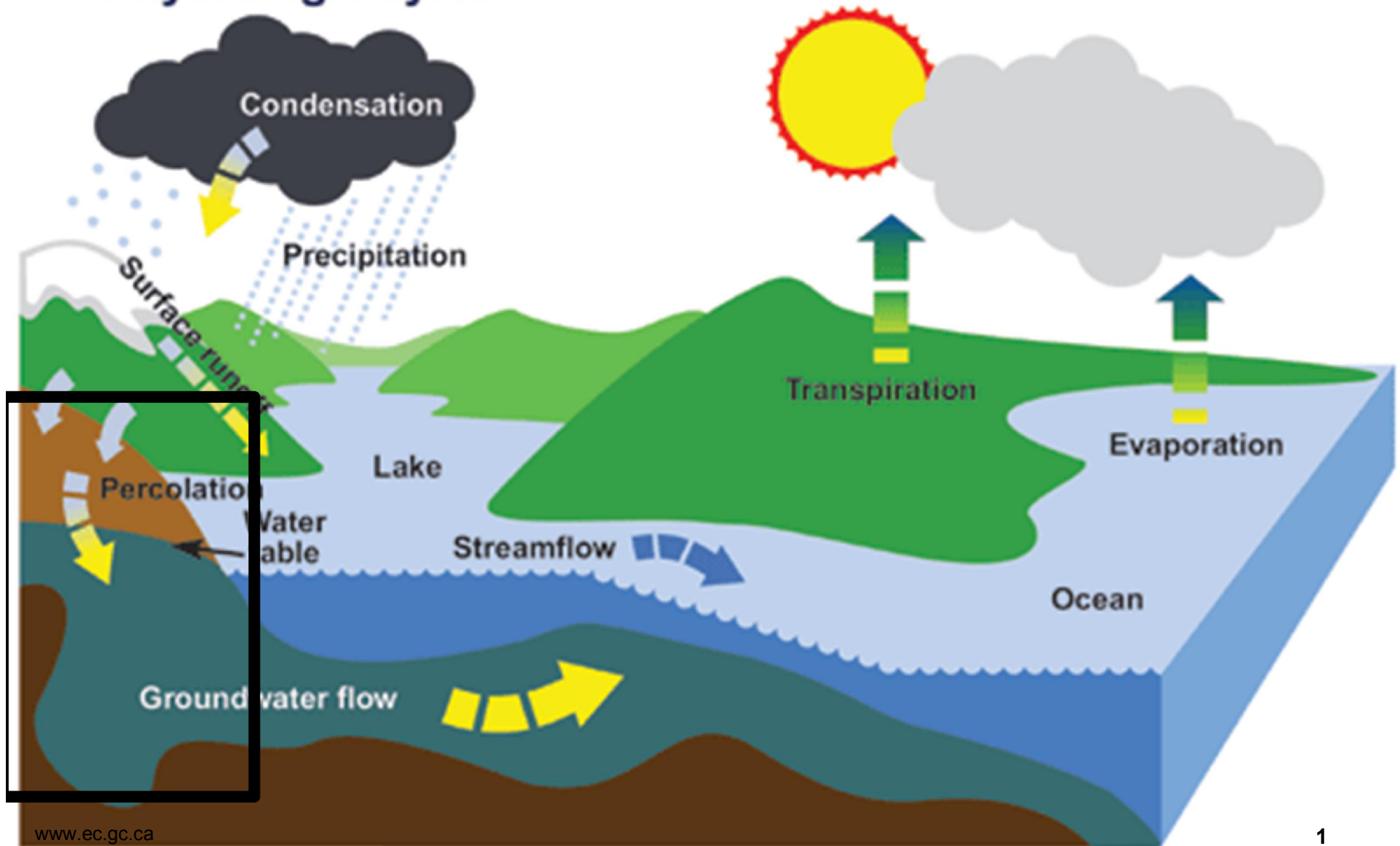


UNIVERSITÉ DE STRASBOURG

SIMULTANEOUS ESTIMATION OF GROUNDWATER RECHARGE AND HYDRODYNAMIC PARAMETERS FOR GROUNDWATER FLOW MODELING

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The hydrologic cycle



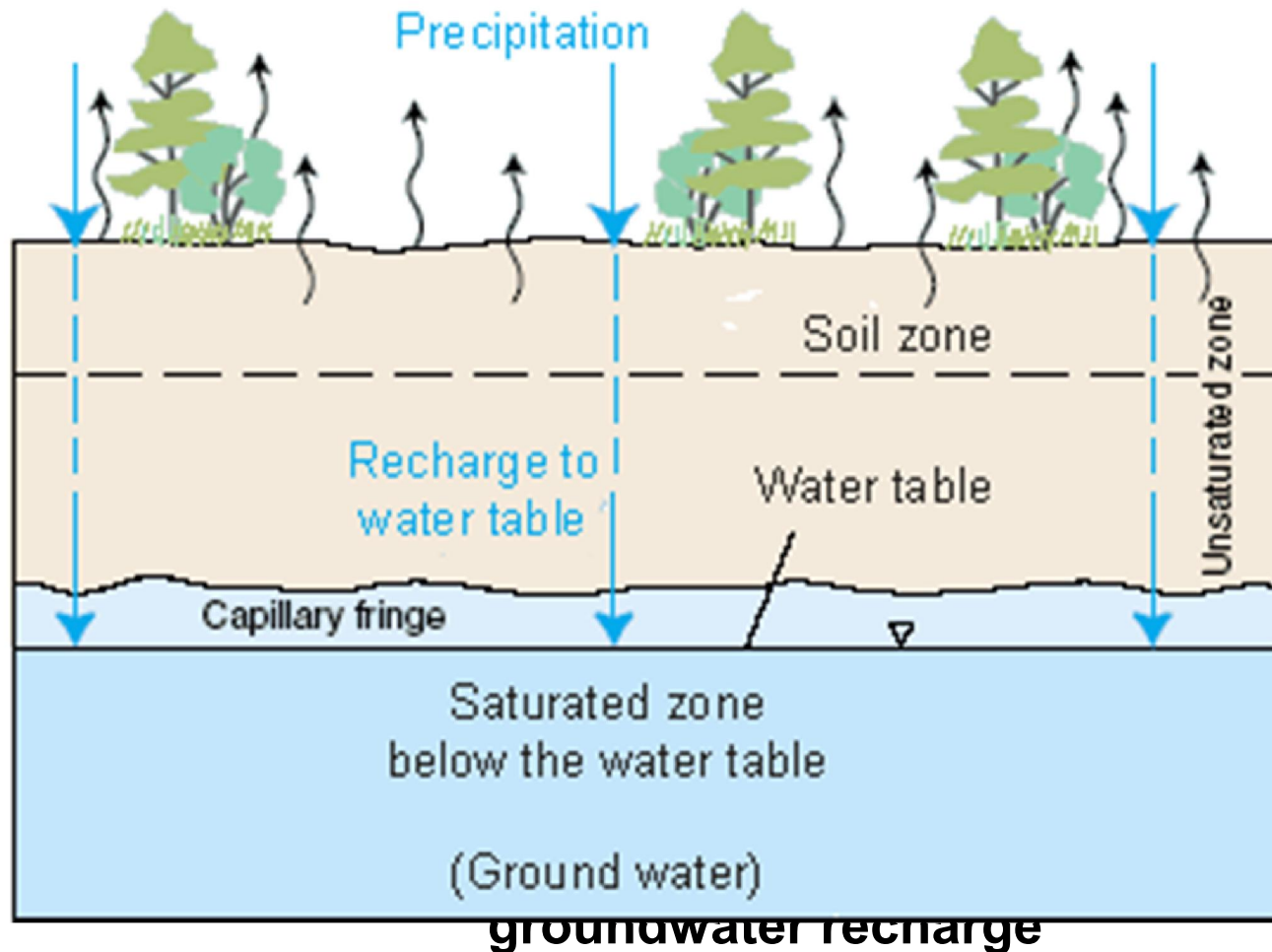
- **Groundwater (aquifers):** 3rd freshwater reservoir on the planet
- Replenished essentially by precipitation through **groundwater recharge**

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s (Hydraulic heads)

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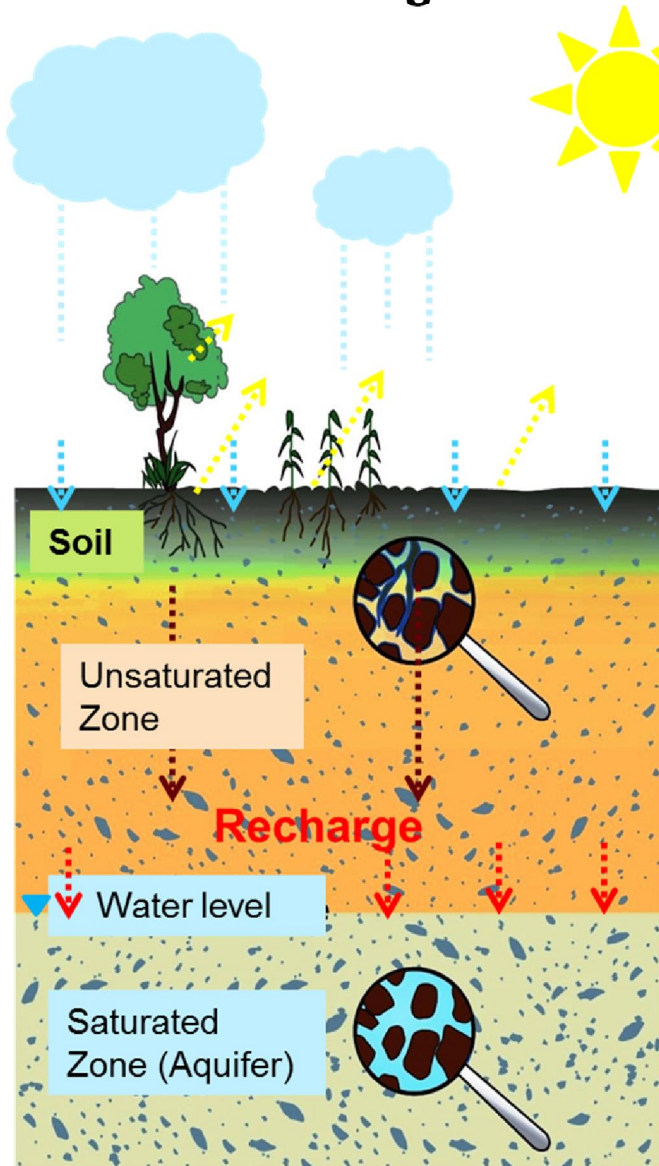
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it, q_s includes

Recharge : water transfer from precipitation to groundwater



- ☺ Groundwater flow models: **Sophisticated and accurate**
- ☹ Recharge modeling: **Challenge**
 - **Complex hydrological component**

Groundwater recharge depends on:

- Climatic conditions
- Vegetation
- Soil and root zone
- Unsaturated zone

Values can not be directly measured



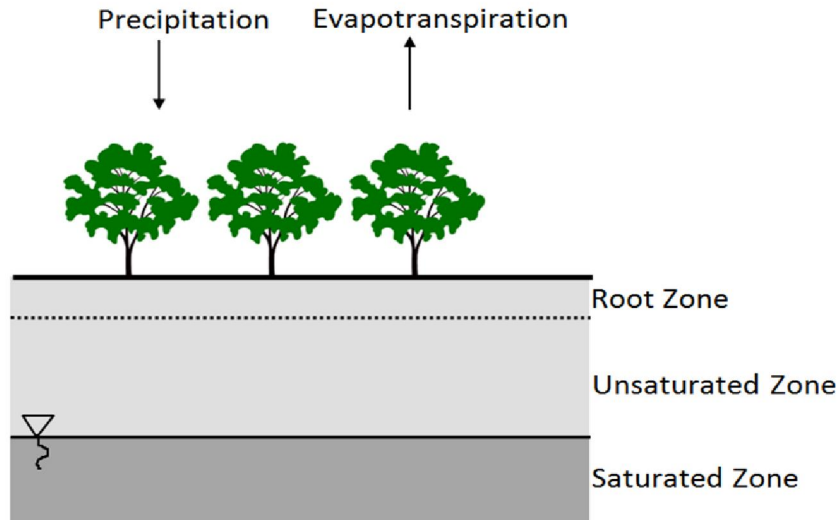
Estimation

- Direct methods: lysimeters, TDR
- Empirical methods
- Tracers
- **Conceptual model**

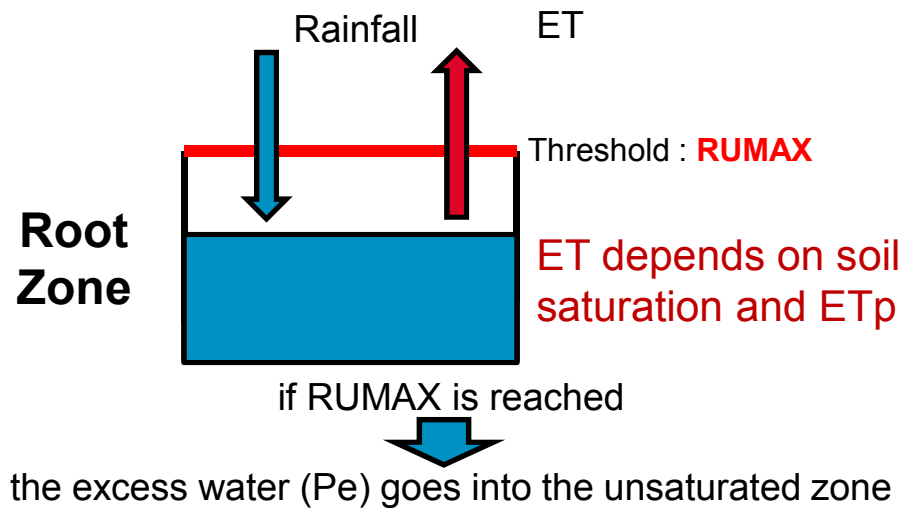
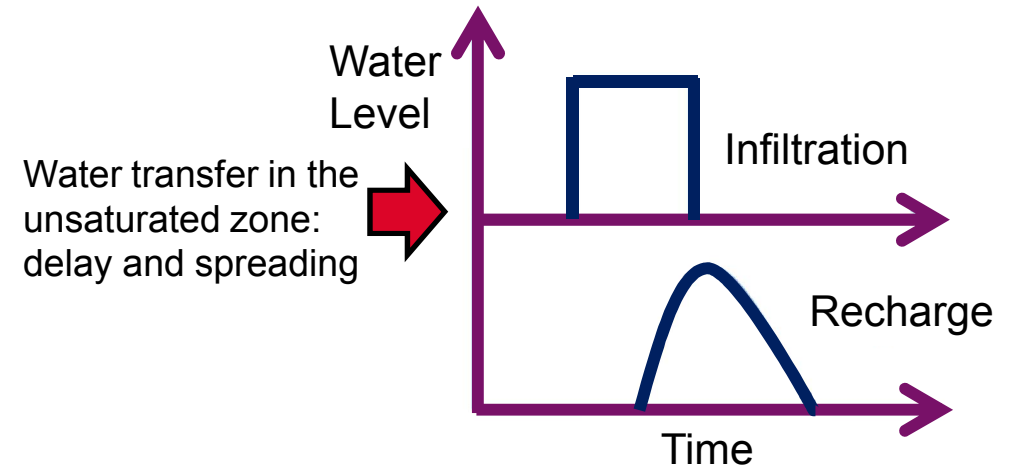
Complexities of recharge



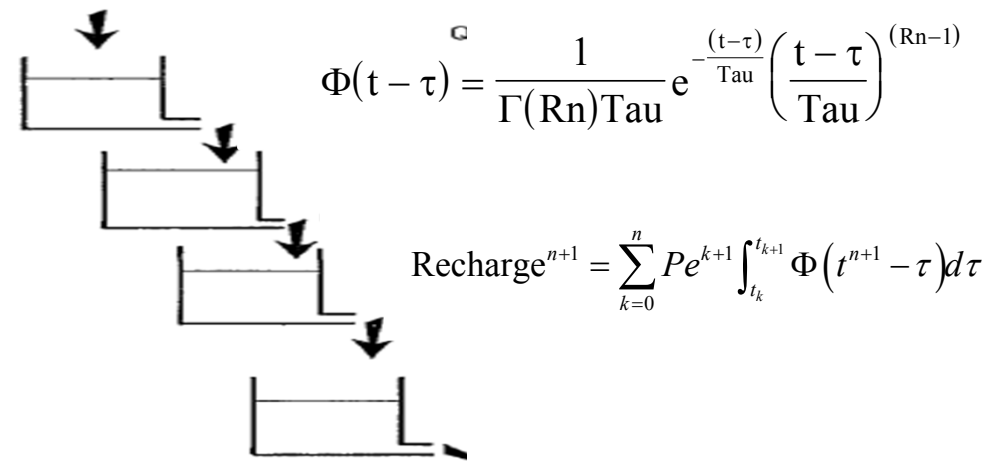
Physical models are difficult to use



Hydrodynamic in the unsaturated zone

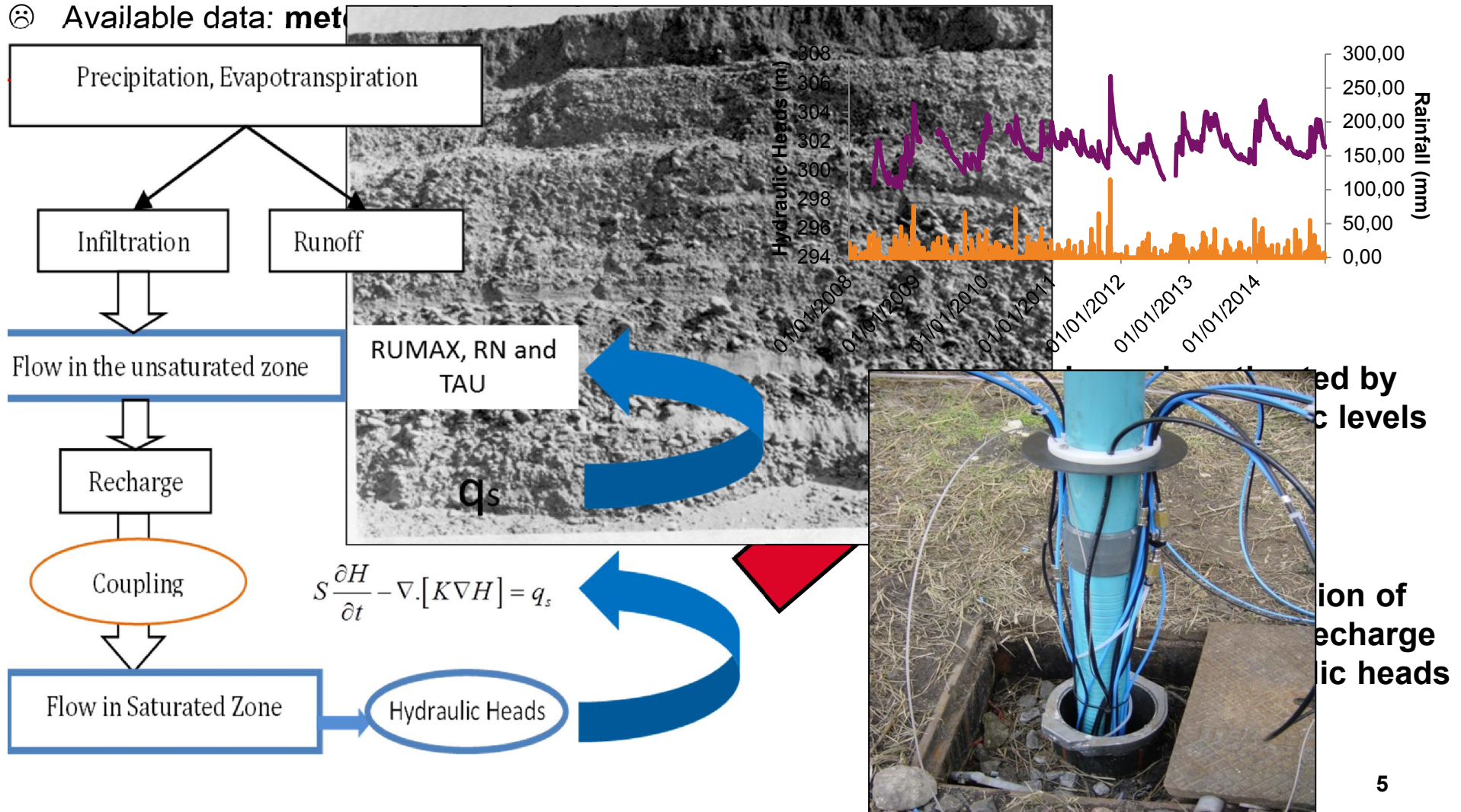


Modeling : propagation of a unit pulse signal



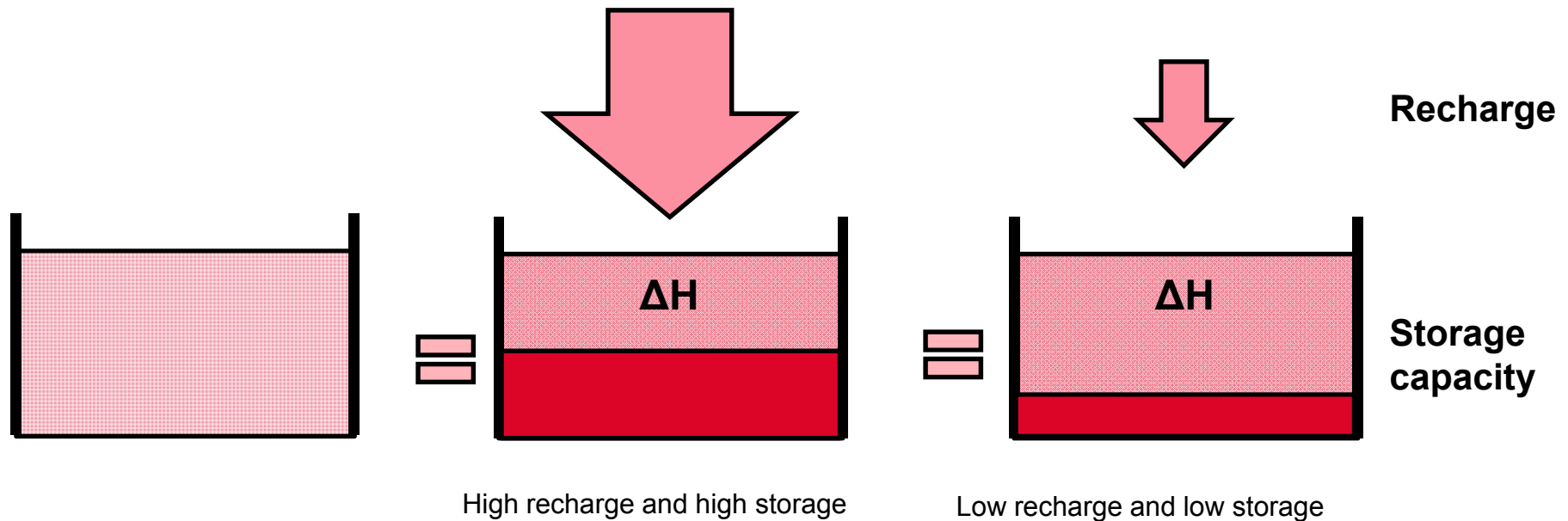
Nash parameters: **RUMAX, RN, TAU**

- ☹ Model uses constants called parameters
- ☹ **Aquifers are highly heterogeneous** → Parameters are not known accurately or even **unknown** in natural environments
- ☹ Available data: **met**



☹ **Correlation between recharge and storage capacity (S and q_s)**

Recharge and storage capacity lead to the same effect



- ☺ Calibration is performed over time and space
- ☺ But... recharge is variable over time while the storage capacity is constant
- ☺ Aquifer has two states : transient and steady

OB24

Slide 7

OB24

voir si on laisse comme ça le KgradH...?

Olivier BILDSTEIN; 07/09/2016

Diffusivity equation

$$S \frac{\partial H}{\partial t} - \nabla \cdot [K \nabla H] = q_s (RUMAX, RN, TAU)$$

➤ **When recharge is zero (summer)**

- Steady state

$$-\nabla \cdot [K \nabla H] = 0 \quad \rightarrow \quad \text{Determination of } K$$

- Transient state

$$S \frac{\partial H}{\partial t} - \nabla \cdot [K \nabla H] = 0 \quad \rightarrow \quad \text{Determination of } S \text{ knowing } K$$

➤ **During recharge period (winter)**

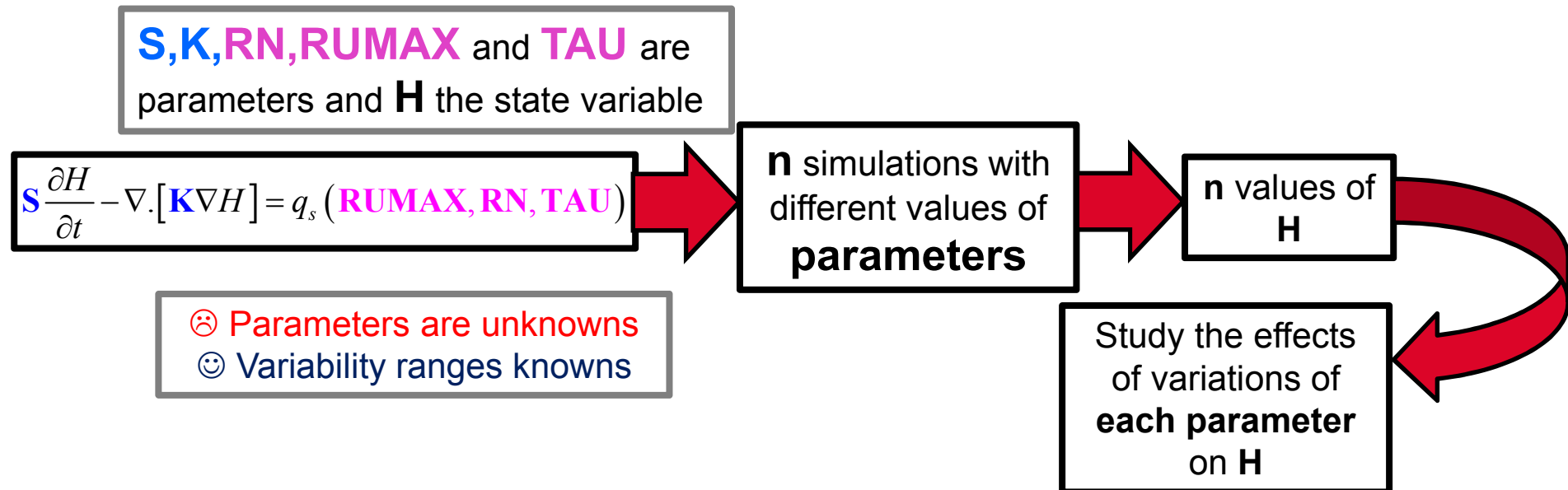
$$S \frac{\partial H}{\partial t} - \nabla \cdot [K \nabla H] = q_s (RUMAX, RN, TAU) \quad \rightarrow \quad \text{Determination of } RUMAX, RN \text{ and } TAU \text{ knowing } K \text{ and } S$$

😊 Global sensitivity analysis

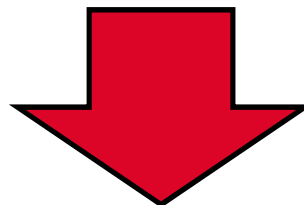
Assess the calibration approach and guide the parameter estimation

Uncertain parameters: variability ranges known → Quantify the effects of parameters uncertainty

Global sensitivity analysis: focuses on the output uncertainty over the entire range of values of the input parameters both single and in combination with one another

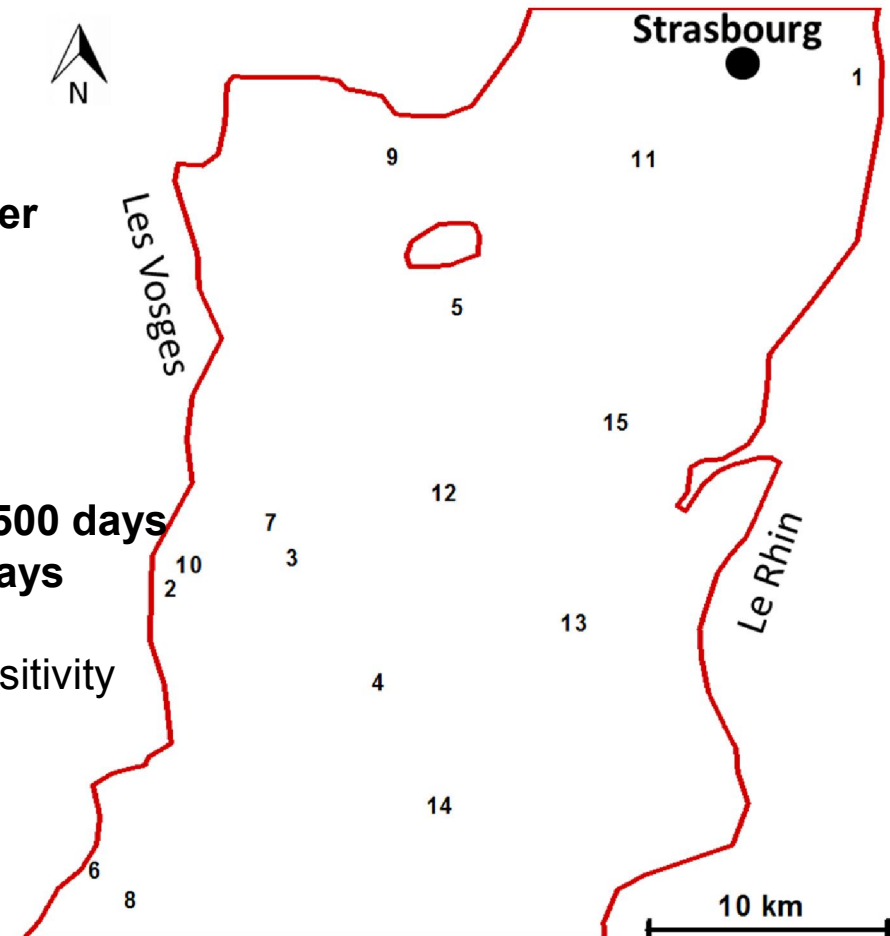



Variance based approach:
Sobol Indices



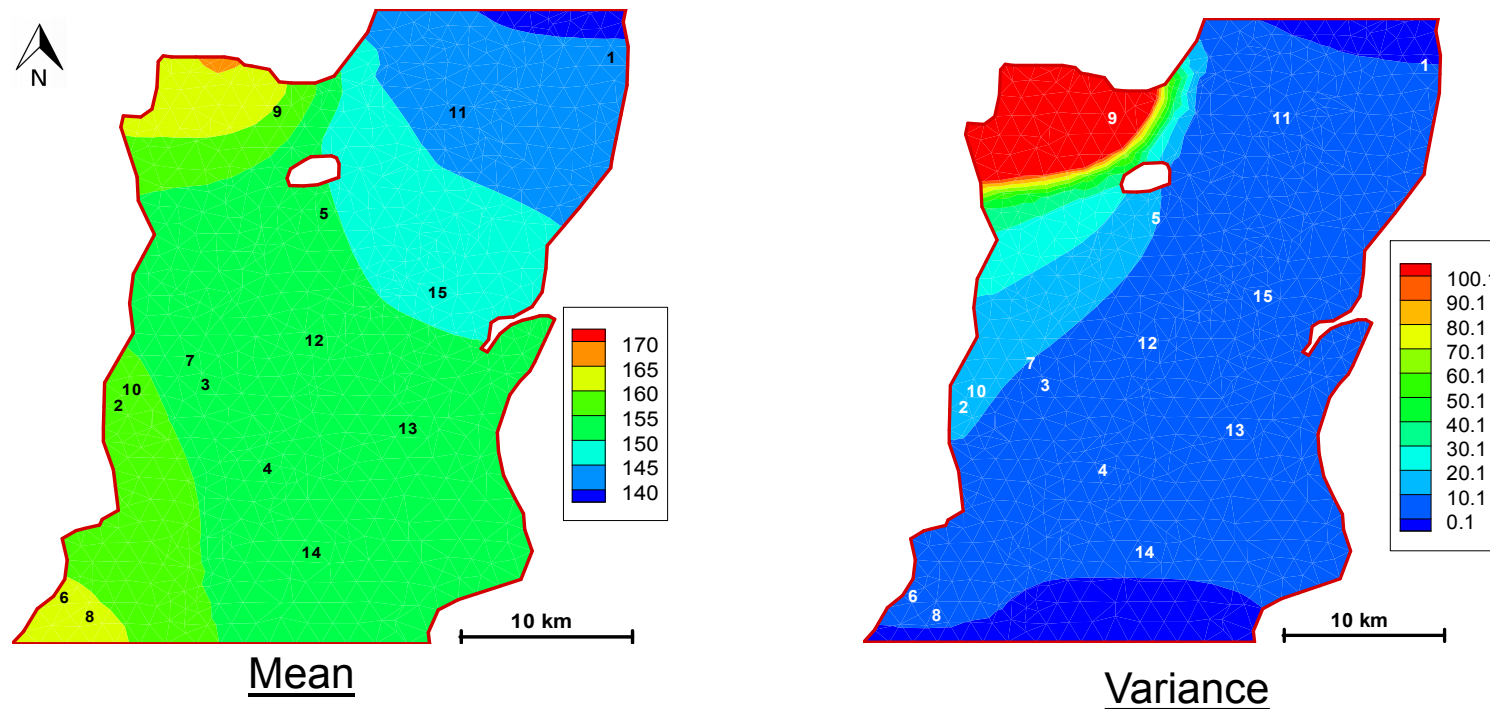
Surrogate model: Chaos polynomial expansion

- **A subdomain of the Upper Rhine alluvial aquifer** widely studied at Université de Strasbourg
- Average rainfall per year: **1100 mm**
- Average of evapotranspiration per year: **900 mm**
- Simulate variations of piezometric levels during **1500 days**
- Sensitivity analysis performed over the last **450 days**
- Investigate spatial and temporal variations of sensitivity indices



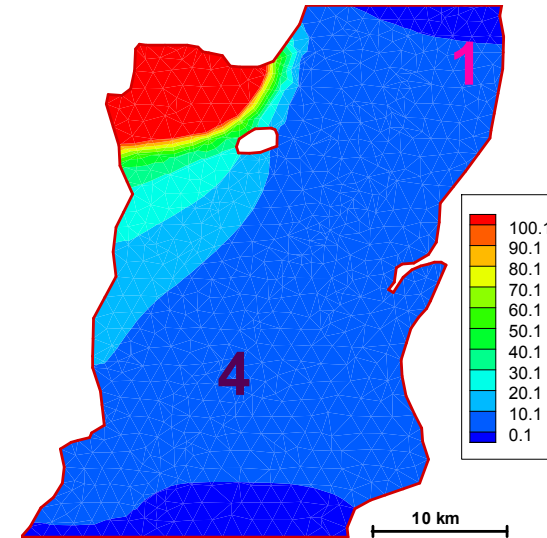
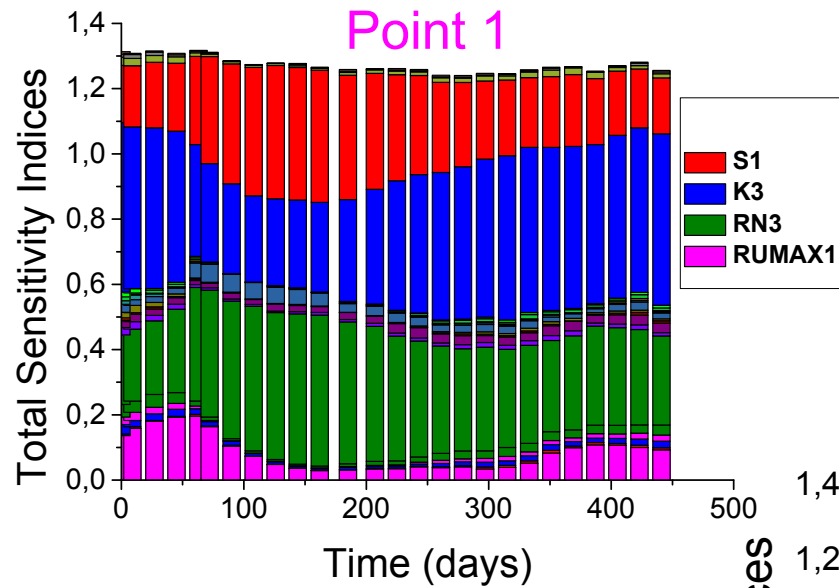
- Nash parameters for recharge: **RUMAX, RN, TAU**
- Aquifer parameters: **K and S**
- Parameters defined by zones (**Zonation**)  20 parameters

Mean and variance of the hydraulic head

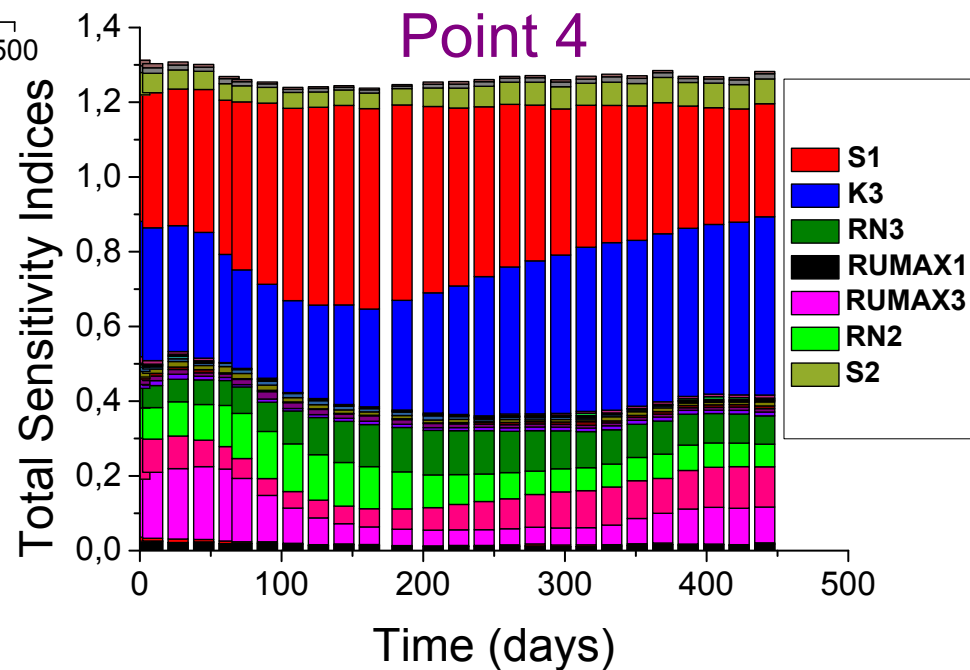


- Hydraulic heads range between 140 and 165 m
- Variance ranges between 0 and 100: the imposed boundaries have the lowest variance

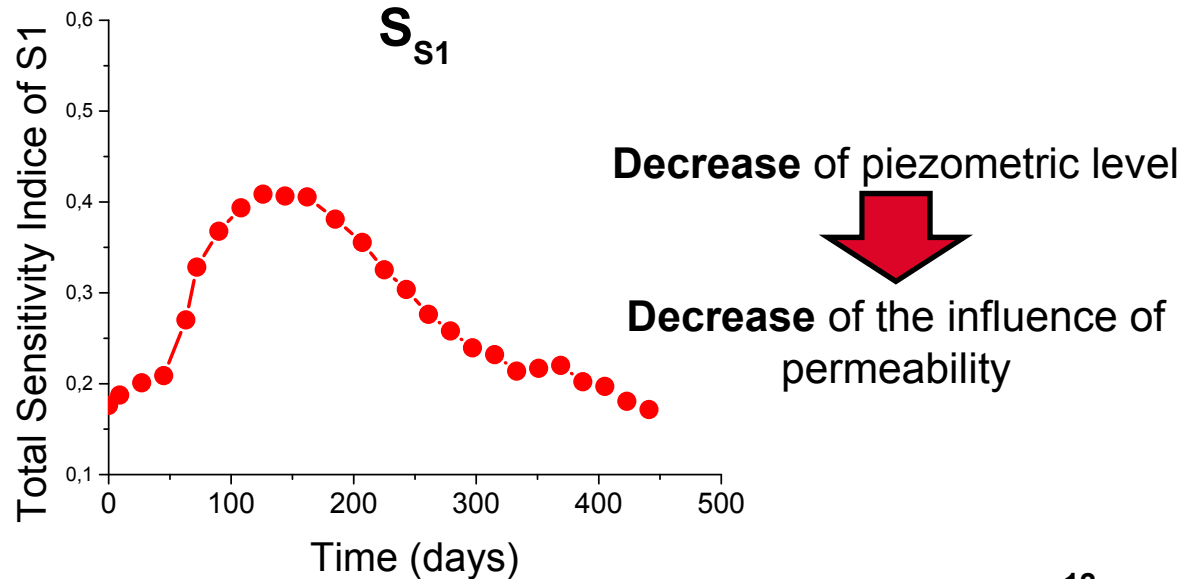
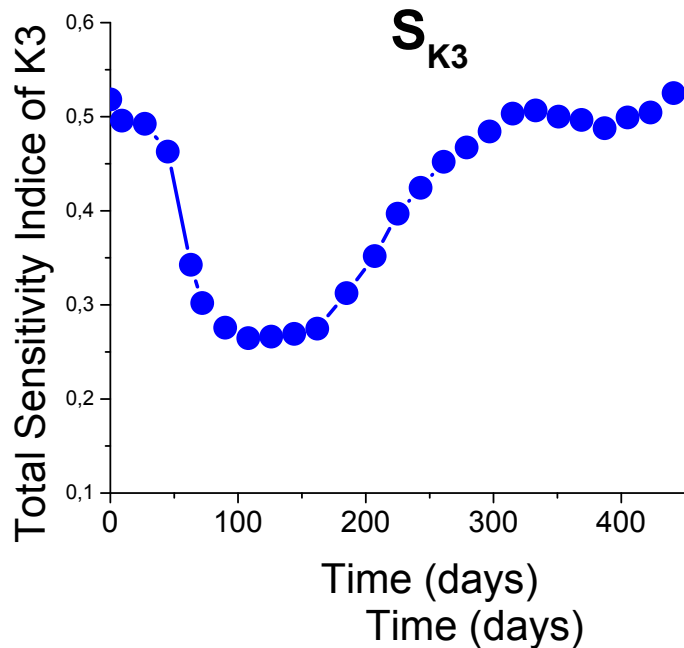
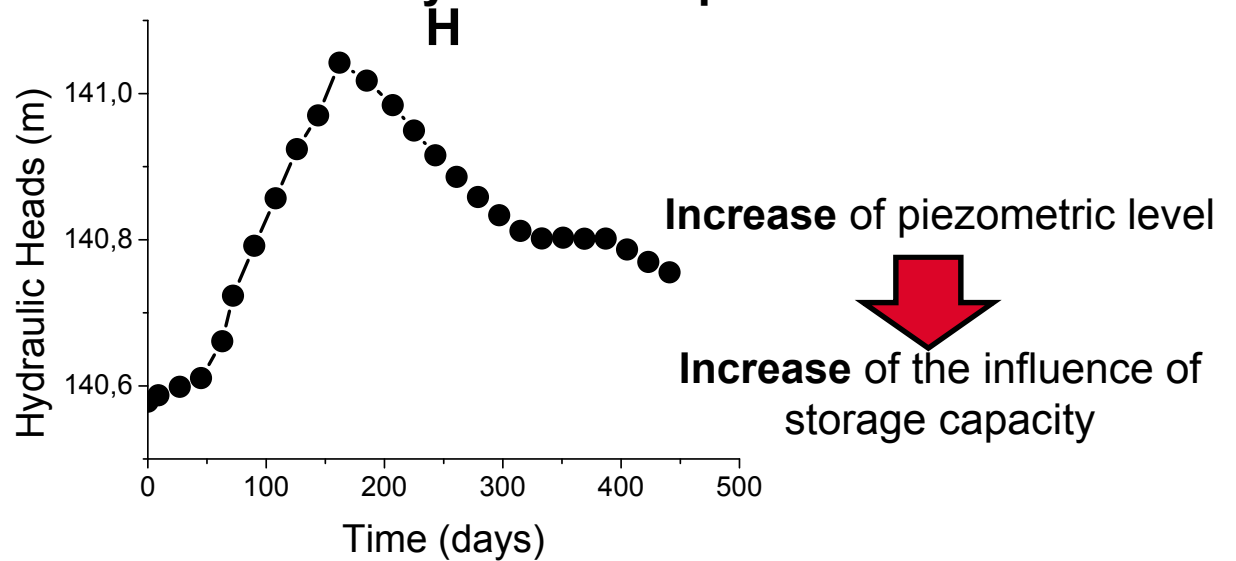
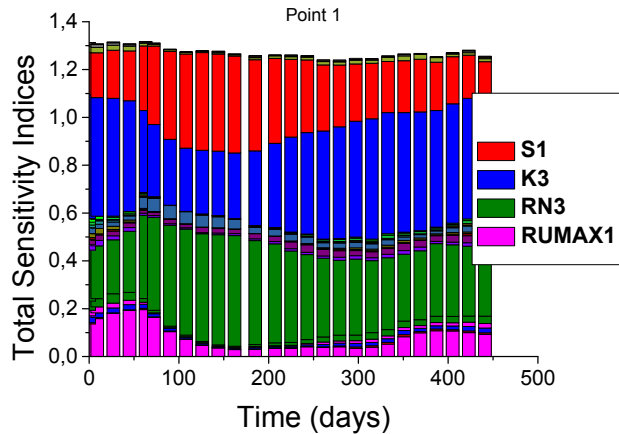
Temporal variation of sensitivity indices



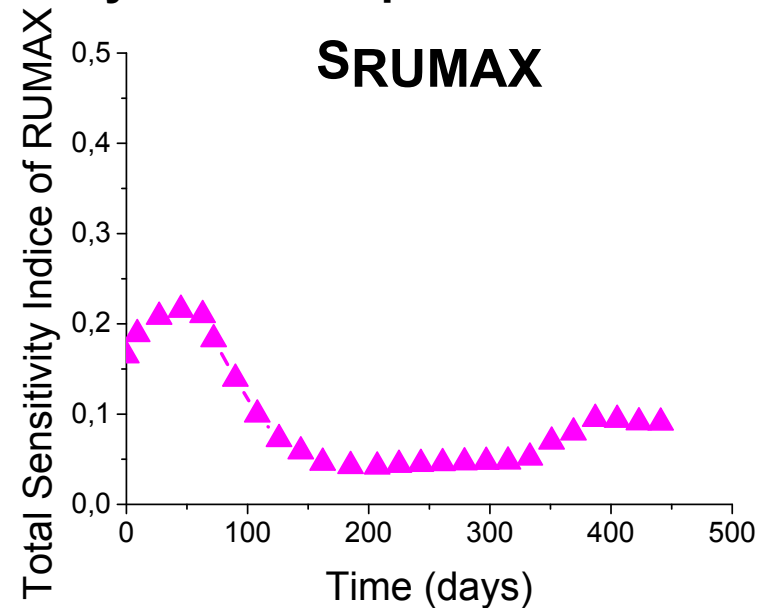
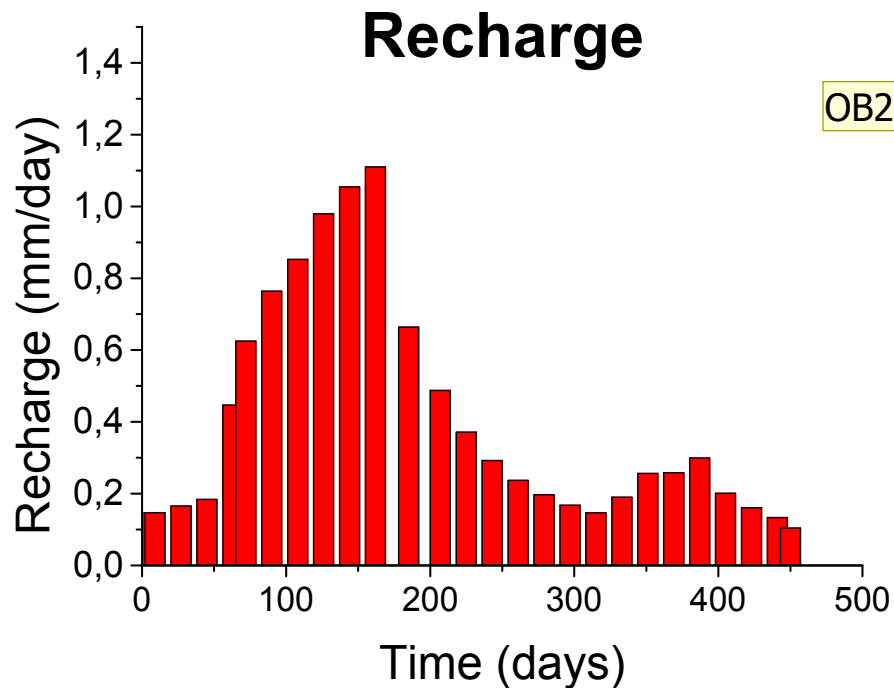
Influential parameters: **S**, **K**,
RUMAX and **RN**



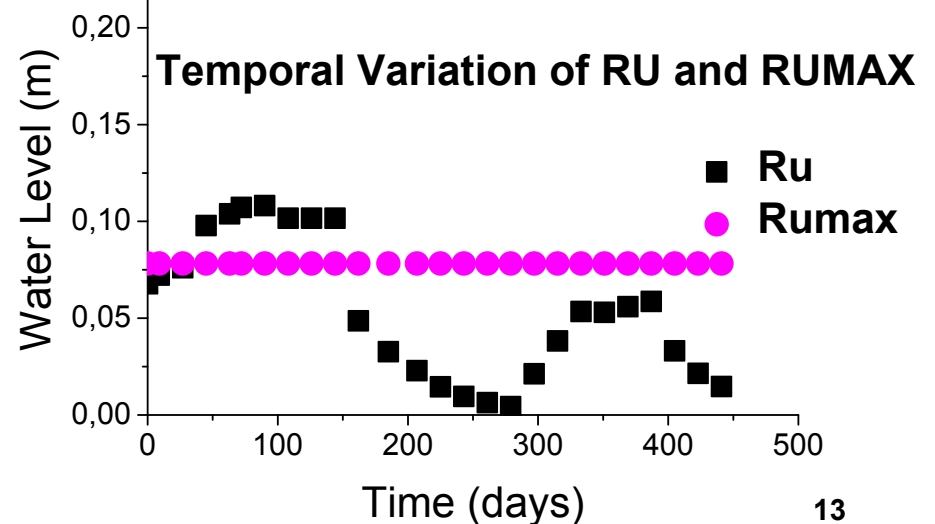
Temporal variation of sensitivity indices at point 1



Temporal variation of sensitivity indices at point 1



RN: water transfer in the unsaturated zone (delay and spreading)
RUMAX: threshold parameter



Slide 14

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attention: enlever les barres d'erreur

Olivier BILDSTEIN; 07/09/2016

➤ Determination of influential parameters

Nash: RUMAX and RN

Groundwater flow: K and S

➤ Interactions between the parameters are not really significant compared to the parameters' main effects

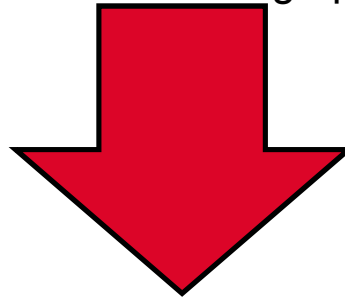
➤ Temporal variation of parameter sensitivity

Nash: saturation of the 1st reservoir (RUMAX) and recharge (RN)

Permeability: decrease of the hydraulic heads

Storage capacity: increase of the hydraulic heads

➤ Weak interactions between aquifer and recharge parameters



Confirmation of the feasibility of simultaneous estimation of recharge and groundwater flow parameters

*Thank you for
your attention*


Let us consider a model with y as output and n input parameters, X the parameters vector

➤ **ANOVA decomposition** : $y = f_0 + \sum_{i=1}^n f_i(x_i) + \sum_{j>1}^n f_{ij}(x_i, x_j) + \dots + f_{1,2,\dots,n}(x_1, x_2, \dots, x_n)$


f_i : contribution of single parameter i to the output

f_{ij} : contribution of the combined effect of parameters i and j to the output

$$H = H_0 + f_S(S) + f_K(K) + f_{RN}(RN) + f_{RN,K}(RN, K) + f_{S,K}(S, K) + \dots$$

Orthogonality of f  $V = \sum_{i=1}^n V_i + \sum_{j>1}^n V_{ij} + \dots + V_{1,\dots,n}$

➤ **Sensitivity (Sobol) Indices** : $S_i = \frac{V_i}{V}$; $S_{ij} = \frac{V_{ij}}{V}$

Computation of f  Polynomial expansion

➤ **Chaos polynomial expansion**

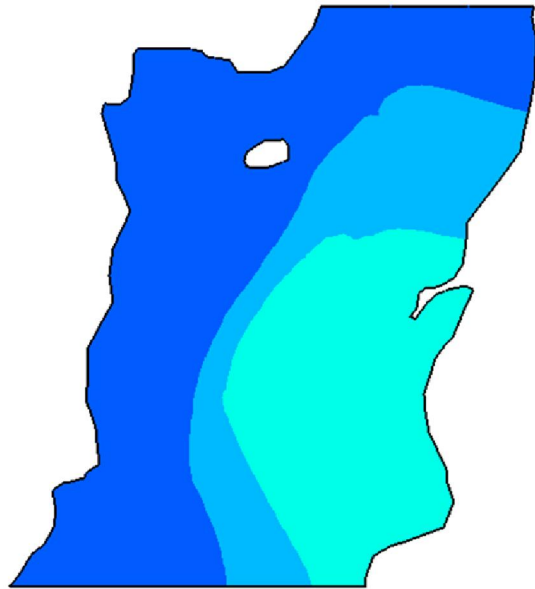
$$y = \sum_{j=0}^{+\infty} a_j \psi_j(x_1, \dots, x_n)$$

x : variable

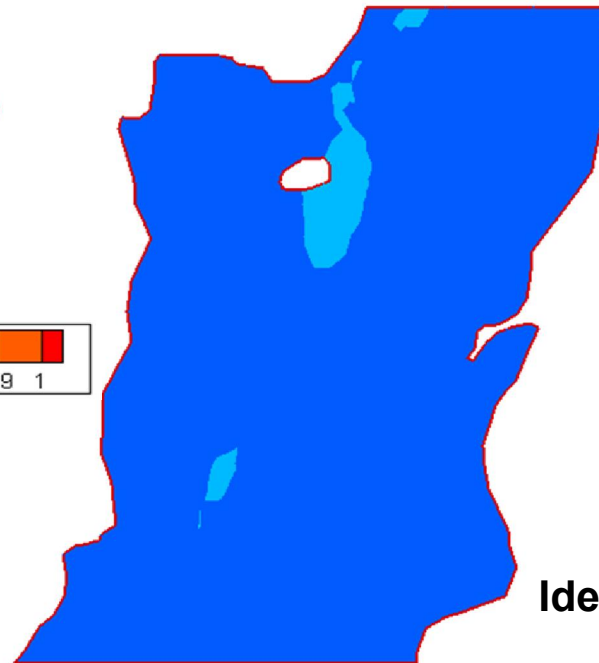
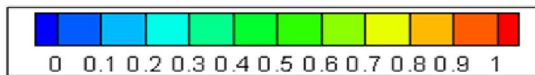
Ψ : orthogonal polynomial (Legendre, Hermite)

a : coefficients determined by least squares regression

Sensitivity maps over time



S1



RUMAX3

RUMAX3: sensitive from the **1st** to **185th** day

S1: influential between the **185th** and the **346th** day



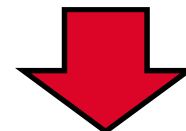
Weak or even negligible correlations between the 185th and the 346th day



Identification of S



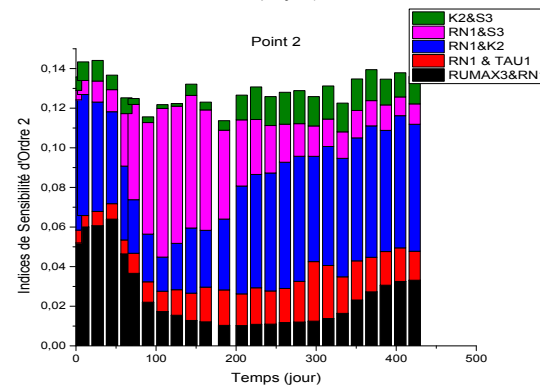
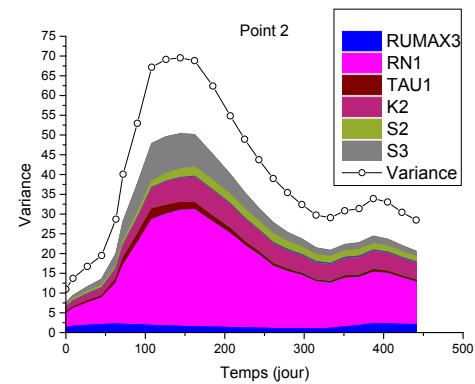
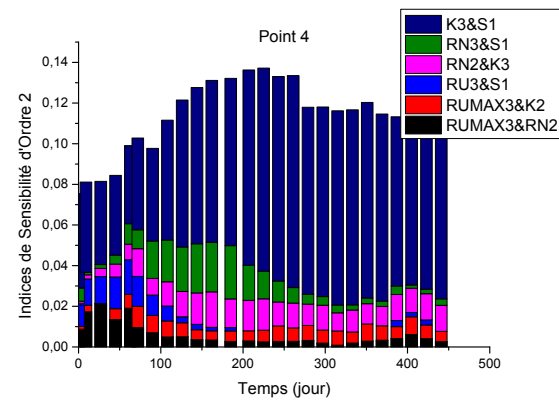
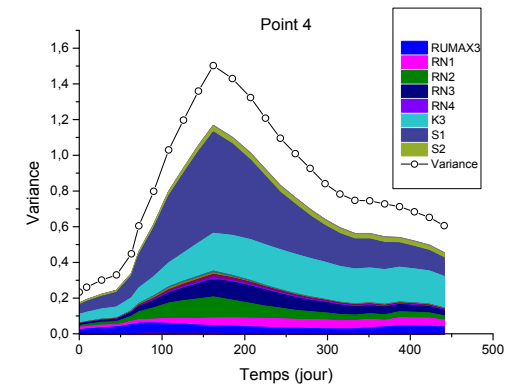
During the first days



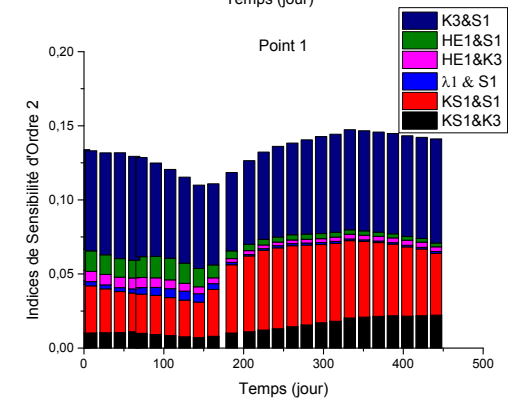
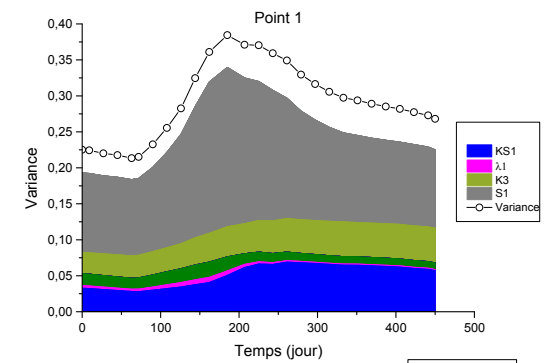
Identification of RUMAX and RN knowing S

Parameters Interactions

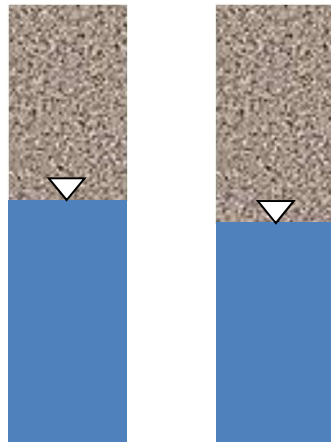
Nash model



Richards model



Assumption of constant hydraulic head over decade



$$q_i^n = -K \frac{\partial}{\partial z} (h - z) = K \left(h_{i+\frac{1}{2}} \right) \left(\frac{h_i^n - h_{i+1}^n}{\Delta z_i} + 1 \right)$$

If the hydraulic heads decreases, the imposed pressure decreases, and the flux calculated is positive and strong

If the hydraulic heads increases, the imposed pressure increases, and the flux calculated is negative

RESOLUTION OF RICHARDS EQUATION

